









The test for determining navigability used in this analysis is from *Defenders of Wildlife v. Hull*, 199 Ariz. 411,426, 18 P.3d 722 (App. 2001):

Also, physical evidence is presented on two issues: (1) navigability or non-navigability of the Verde River in its "ordinary and natural condition" at the State of Arizona's admission to the United States on February 14, 1912, consistent with the Arizona Court of Appeals decision in *State v. Arizona Navigable Stream Adjudication Comm'n*, 224 Ariz. 230, 229 P.3d 242 (App. 2010); and (2) segmentation of the Gila River consistent with the United States Supreme Court's decision in *PPL Montana, LLC v. Montana*, 556 U.S. ____, 132 S.Ct. 1215 (2012).

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View looking upstream at Verde River Channel from below Black bridge in Camp Verde. The river channel has scoured through the alluvial sediments (Holocene material) into the Verde Formation (light colored material). Hjalmarson





























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There obviously is a considerable amount of silt, sand and gravel at

this location.

Coalition, 179p.

Photo from Evans, K and McClain, C., 2005, Wild and Scenic River Proposal for The

Upper Verde River, In conjunction with the Arizona Wilderness

Bedrock bank

Photo from: Bowman, S. N., 2001, VERDE RIVER TMDL

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FOR TURBIDITY, Arizona Department of Environmental

Quality, 33p.



















Total cultivated land Location Acres Flow, cfs 1 Granite, Williamson Valley, Walnut, and Big Chino Creeks 8095 35 USGS Clarkdale gage 8215 36 Base flow lost from Verde River because of diversions for irrigation of cultivated land. Diversions typically are low dams and shallow wells in stream sediment and cultivated land typically is on Holocene sediments (Lynx soil series that is recent alluvium (Wendt, 1976). mount of base runoff lost to ET from cultivated land shown in column 4 (2of2). Wendt, G. E. and others, 1976, Soil survey of Yavapai County, AZ -Western Par U. S. Soil Conservation Service 1210. August 2 -Western Par	mputation of total base runoff		METHOD			
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METHOD 2

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The irrigated land for these areas is given in the Hayden (1940) report. The approximate 100 cfs loss to ET was simply distributed between the two areas on the basis of the ratio of irrigated acres for the two areas. About 28 cfs was lost to ET from irrigated land above gage 09503700 and 72 cfs was lost in the Verde Valley.

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METHOD 3 Method 3 Conveyance-slope estimates of historic base runoff using Federal Land Survey Estimates of base flow in Verde River at east side of section 12, T17N R2W on May 1909

and at boundary between sections 1 and 12, T17N R1W also during May 1909 using width and depth of Federal Land Surveys.

data.

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MIDDLE-LOWER REACH

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Important parts of this analysis of the hydrology below USGS gage 09504000 include:

- 1. USGS records of stream flow at gages 09503700, 09504000, 0950600 and 09510000
- 2. A report by the USBR (1952) that calculated the Virgin flow for the mouth of the Verde River
- A report by the USGS (HA-664 by Freethey and Anderson (1986)) that estimated base runoff (Qbfa, the 90th percentile of daily discharge) for the basin fill and underlying aquifers.

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MIDDLE-LOWER REACH

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•Base runoff of the lower reach of the Verde River is simply computed by adding (1) the 100 cfs difference between the Virgin average annual runoff and the gaged average annual runoff at gage 09510000 (recall method 2 for upper reach) that was associated with early settler use of base flow and (2) the base flow (from USGS HA664) at and below the USGS gage 09506000.

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MIDDLE-LOWER REACH ·Base runoff of the lower reach of the Verde River is simply computed by adding (1) the 100 cfs difference between the Virgin average annual runoff and the gaged average annual runoff at gage 09510000 (recall method 2 for upper reach) that was associated with early settler use of base flow and (2) the base flow (from USGS HA664) at and below the USGS gage 09506000. •The base runoff associated with Qqa and Qmf was simply distributed across the middle Verde River between USGS gages 09504000 and 09506000. • The resulting natural base flow is shown in on the following slide. Hjalmarson for ANSAC 69







A summary of a detailed assessment of the 36.6 mile reach of the upper Verde River from the dam at Sullivan Lake to the USGS stream gage near Clarkdale, AZ is presented first and is followed by a summary of a more general assessment below the Clarkdale gage to the mouth at the Salt River.

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Large debris (boulders) from side slopes. Obviously only very large high (kinetic) energy flow will move such large obstructions. Also, energy is lost (with a corresponding decrease of velocity and increase in depth) as streamflow encounters this rough channel material





Photo of cobbles near Sycamore Canyon .(Photo by James Cowlin USFS).

The transport of sediment debris by rivers like the Verde River is common knowledge.

The forces (eg.-shear forces) involved in shaping and maintaining the channel are related to both the amount and duration of water flow. As flow (energy) in this scene increases, the silt and sand can become suspended in the flow and the gravel, cobbles and small boulders can be moved by pushing, rolling and skipping. The rate of sediment transport is much less for base flow than floodflow but the duration of base flow is considerably longer. Hjalmarson for ANSAC 78

Many cross sections with channel widths have been measured by the Federal Surveyors, U S Forest Service, Sierra Club and the USGS upstream of the USGS Clarkdale gage.
Many current meter measurements have been made by

•Many current meter measurements have been made by the Sierra Club, USFS and USGS along the river upstream of the USGS gage near Clarkdale.

•Downstream of the Clarkdale gage many cross sections were measured by the Federal Surveyors, Finally, the USGS operates stream gages 09503700, 09504000, 09506000, 09508500, 09510000 and 09511300 where many current meter measurements and a few rating curves are available.

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USGS data for June 13 and 14, 2000	June 2000
Miles Site name (approx.)	Discharge Width cfs ft
0 Sullivan Dam	
9.8 Verde River near Paulden	21.0 14.8
11.5 VERDE RIVER AT BULL BASIN CANYON	19.0 18.7
13.5 VERDE RIVER ABOVE DUFF SPRING	20.0 26.8
14.4 VERDE RIVER BELOW DUFF SPRING 2	23.0 29.2
17.8 VERDE RIVER ABOVE HELL CANYON	19.0 26.6
18.2 VERDE RIVER BELOW HELL CANYON	17.0 50.0
19.4 VERDE RIVER AT US MINE 2	17.0 15.9
23.7 VERDE RIVER ABOVE PERKINSVILLE DIV.	* *
24 VERDE RIVER NR PERKINSVILLE	15.0 31.0
26 VERDE RIVER BELOW ORCHARD FAULT	* *
28 VERDE RIVER ABV MORMON POCKET	26.0 41.4
32 VERDE RIVER NEAR BM 1813 (abv Syc. Ck) 58.0 44.2

















Location	Mean annual		Median		Q90			
	QIV	ax. Depth	0	Max. Depth	Q M	ax. Depth		
	CIS	n	ers	n	crs	n		
mile 0	00	27	60	2.4	E4	2.2		
mile 3.3	80	2.7	60	2.4	54	2.3		
Sip	00	2.9	00	2.0	54	2.5		
mii 6.8	80	4.4	60	3.9	54	3.8		
Paulden	80	2.8	60	2.4	54	2.4		
mile 16 Reer Siding#	80	3.8	60	3.3	54	3.2		
bear Siding"	00	3.4	00	3.1	54	3.0		
mile 23.3 Derkine ville	80	4.4	60	4.2	54	4.1		
reikinsville	80	2.0	60	2.5	54	2.5		
mile 25	100	2.2	100	1.9	04	1.9		
mile 32.2	190	4.2	100	3.0	94	2.9		

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•Most of the Verde River is pools where riffles occupy a much smaller portion of the river.

•Thus, typical depths for natural conditions along the reach from mile 3.3 downstream to the USGS Clarkdale gage are at least 3.5 ft (mean annual), more than 3.0 ft (median, Q50) and about 3 ft. (Q90). •Also, the depths closely represent depths along a potential navigation lane (or corridor) used for small water craft.

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•There are alternating pools and riffles along the Verde River and many of the riffles are located at the mouths of tributaries that dump flood debris into the Verde River.

•Most of the channel bed is gravel and cobbles with sand and boulders.

•Most of the conditions along the Verde River are typical of many perennial gravel bed streams and streams where the bed material is larger than coarse sand.

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•The evidence in my report shows the Verde River upstream of Horseshoe reservoir behaves like a typical channel where "A natural channel migrates laterally by erosion of one bank, maintaining on the average a constant channel cross section by deposition on the opposite bank.

•In other words, there is general equilibrium between erosion and deposition.

•The form of the cross section is stable, meaning more or less constant, but the position of the channel is not." (As described by Leopold, 1994, p.5 for rivers in general.).

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The evidence also suggests there was a single well defined main channel along the entire river following the large floods of 1891 and even 1993.
Downstream of Horseshoe Dam where nearly

all flow is from controlled releases from the two major reservoirs, the main channel typically is well defined but there are a few braided reaches where the recent channel(s) of the lower Verde River is (are) not considered representative of the natural condition.

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Thus, while much of the size and shape of the natural main channel are considered approximately the same as the recent channel for this study, it is likely that flow in the recent channel with the highly regulated flow below Horseshoe Dam is shallower and appears wider than was the natural channel.

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Evidence shows active stream channel deposits composed of very poorly-sorted sand, pebbles, and cobbles with some boulders to moderately-sorted sand and pebbles.
Channels are generally incised 3 to 7 ft below adjacent recent terraces. Channel morphologies generally consist of a single thread high flow channel or, in places, multithreaded low flow channels with gravel bars.
These active channels of recent silt to boulder material convey base flow, direct runoff and flood flow.
Downstream of Bartlett Dam where river sediments have been intercepted and stored the channel banks have eroded with some channel braiding.







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- There was excellent agreement among the three independent estimates of natural runoff to the upper Verde River.
- These techniques use published information of the USBR, USGS, USFS, Salt River Project, local historic newspapers and Federal Land Surveys.
- Also, surveyed channel widths of the original land surveys, that were considerably greater than recent measured widths, support the estimated amount of natural runoff.
- Base runoff along the entire river conforms to the amount of virgin flow (USBR, 1952) at the mouth.

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Published information of th Standard engineering methods narkable fundamental Station Q90 Median (Q50) Mean annual ydrologic/morphologic principles cfs 60 116 cfs 80 211 54 111 5037000 9504000 stematic three-step method (hydrology, hydraulics This completes the Hydrology rphology, navigability USGS June : 80 60 40 harson for ANS.



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(Hyra, R., 1978, Methods of assessing instream flows for recreation: Instream Flow Information Paper No. 6, U. S. Fish and Wildlife Service and others, 14p.)

Hjalmar BORATOTANSAC 12/8/2014



•The channel has shifted in plan view but the shape and size of the main channel have not changed much.

•The series of pools (deep water areas typically behind riffles) and riffles (shallow water areas typically dominated by cobbles and small boulders) are relatively stable throughout most of the Verde River.

•Downstream of Bartlett Dam the once single channel has become braided in places and the main channel may have become wider and shallower in places.

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